

Summary report on MEDEX studies and scientific results on Mediterranean cyclones causing high impact weather

Andrea Buzzi, Evelyne Richard, Romualdo Romero (Nov. 2005)

1. The physical mechanisms, as derived from theoretical, observational and numerical studies.

- Mediterranean cyclogenesis is often initiated as secondary cyclogenesis. This is true especially for orographic cyclogenesis (Speranza, 2001; Buzzi et al, 2003). Orographic cyclogenesis is particularly prominent in climatology (since 1950's studies, e.g. Petterssen, 1956). Orographic and especially Alpine cyclogenesis (or cyclogenesis in the Gulf of Genoa) studies initiated in the 50-60's of the last century (for example, Radinovic, 1965; Egger, 1972; Buzzi and Tibaldi, 1978) and culminated in the WMO-ALPEX programme research. It is related to many forecasting problems and high impact phenomena, as for example storm surges (Zampato et al, 2005), Mistral and Bora (Horvath et al, 2005), precipitation especially in Central-Eastern Mediterranean and Eastern Europe. Orographic cyclogenesis north of the Algerian coast, under southwesterly upper-level flow over the Atlas mountains, produces most of the times weak cyclones but that can be crucial for driving flash-flood situations in eastern Spain (Homar et al. 1999; Romero et al. 2000). A well-known area of cyclogenesis is that found in the southeast of the Atlas mountains in Algeria. It has been shown (Prezerakos, 1990; Prezerakos et al., 1990; 2005) that the so-called Saharan or Atlas depressions deriving from this cyclogenesis and migrating usually over the Mediterranean as vigorous cyclones are baroclinic disturbances developing when the polar jet stream moves much further south from its normal seasonal position and approaches the subtropical jet stream. Orographic cyclogenesis affects also the area south-east of the Pyrenees, the Adriatic Sea, the Cyprus area south of the Anatolian mountains.
- Different physical “ingredients” and “factors” have been found to be favourable for cyclogenesis (or cyclolysis): orography (usually cyclogenetic, but depending on the position of cyclones and orientation of main flow and temperature gradient), sea surface temperature, heat fluxes, latent heat release (generally cyclogenetic if coupled with positive heat fluxes), convection, degree of baroclinicity, subsynoptic structure and intensity of the upper-level precursor trough, etc. (Alpert et al, 1995; Prezerakos et al, 1997; Romero et al, 1997; Buzzi et al, 1998; Homar et al. 2002; Lionello et al, 2003; Kotroni et al, 1999; Tripoli et al. 2005; Romero et al. 2005;).
- The importance of pre-conditioning of the Mediterranean (and surrounding areas, e.g. North Africa) atmospheric state prior to cyclone formation, in terms of presence of moisture and low static stability has been documented. Preceding tropical cyclones in the Atlantic and injection of moisture from the Atlantic, the Mediterranean and even from the tropics have been found to determine the water vapour budget and the water cycle that affect the growth rate and the structure of cyclones (Reale et al., 2001; Pinto et al, 2002; Turato et al, 2004; Bertò et al, 2004; Krichak et al, 2004).
- From a theoretical point of view, stability theories have been applied for understanding cyclogenesis: baroclinic instability in the presence of orography; convective and absolute instability (in relation to mobility of disturbances in the growing stage – Fantini and Davolio, 2002); theory of vortex propagation for the cut-off mature stage; symmetric

instability in relation to rainbands (Dobrochotov et al, 2002; Mantovani and Speranza, 2002; Speranza, 2005; Tartaglione and Speranza, 2005).

- Mediterranean small hurricanes ("Medicanes") affect the Mediterranean from time to time as documented mainly, but not only, in satellite images. A couple of such hurricane-like cyclones in the Mediterranean are described in the book by Rasmussen and Turner (2004, pp. 214-219) and compared with polar lows. The dynamics and physics of "Medicanes" are not yet accurately diagnosed due to data sparsity, but have been simulated in several studies (Lagouvardos et al, 1999; Reale and Atlas, 2001; Pytharoulis et al, 2000). Numerical simulations of these systems help to discriminate the role of precursor upper-level cold disturbances and air-sea interaction processes (Homar et al. 2003).

2. Synoptic and subsynoptic-mesoscale aspects and life cycle (genesis, evolution, decay)

- The evolution of fronts associated with cyclones in the Mediterranean is peculiar, especially for orographic or orographically-modified cyclogenesis, due to direct and indirect orographic forcing and dishomogeneity of surface heating and roughness. The presence of (multiple) rainband structures can produce strong orographic and convective precipitation when intersecting orography: Benzi et al., 1997; Buzzi et al, 1998; others...). The "conveyor belts" that in Mediterranean cyclones present similarities and differences from those characterizing oceanic cyclones (Ziv et al, 2005).
- In the context of "PV thinking" diagnostics, the role of upper level PV anomalies and streamers has been demonstrated for Mediterranean cyclogenesis (Prezerakos et al, 1997; Homar and Stensrud, 2004; Kotroni and Lagouvardos, 2000; Lambert et al, 2003; Lambert et al, 2004; Romero, 2001; Romero et al, 2005). Low level PV anomalies and PV banners are generated by the orography and heating and can affect the cyclone life and intensity. Heavy rain events in the Mediterranean basin are often associated with an upstream upper-level trough, generally propagating eastward. For example, in the case of the Algiers flood of 10 November 2001, a deep upper-tropospheric geopotential trough was present from Scandinavia to Morocco, forcing significantly high potential vorticity values down to 700 hPa. Tripoli et al. (2005) investigated the mechanisms leading to the coupling of this upper-level PV anomaly with surface air warmed by wind induced surface heat exchanges (WHISHE). Romero et al. (2004) isolated the mutual interactions of upper-level and lower-level PV anomalies (including diabatically generated PV within the cyclone cloud system) and the background flow. Homar and Stensrud (2004), using the MM5 adjoint system, showed that the structures which influence mostly the baroclinic development were related with the surface front but also with some subsynoptic details of the upper-level anomaly. Lambert et al. (2004), in a series of sensitivity studies to the initial conditions, examined the impact of modifying the geometry and/or the amplitude of coherent PV structures on the localisation of precipitation.
- The role of atmospheric "origin" regions and surface "source areas" of water vapour, especially in cases of heavy precipitation, have been investigated (Turato et al, 2004; Bertò et al, 2004; Krichak et al, 1998; Krichak et al, 2004; Pinto et al, 2002; Lebeauvin et al, 2005). Important origin regions are the Atlantic (at different latitudes, from subtropics to mid-high) but also North Africa and the Mediterranean itself.
- There are different regions of formation and subsequent prevailing paths of cyclones in both the Western and Eastern Mediterranean: Gulf of Lion, Gulf of Genoa, Alboran sea,

Thyrrhenian sea, Northern and Southern Adriatic; Ionian and Aegean seas; Cyprus area (Michaelides et al, 2004; Nicolaidis et al, 2004; Prezerakos et al, 2005); north Africa especially in the lee of the Atlas, etc. (Prezerakos et al, 1999; Nicolaidis et al, 2005). Particular attention has been devoted recently to establish relationships between cyclones and high impact weather, especially heavy precipitation, both over large basins (e.g. Jansa et al. 2001; Gil et al, 2003; Campins et al, 2005a) and for particular regions (e.g. Romero et al. 1999). Systematic intercomparison was made of intense cyclogenesis events affecting the whole of the Mediterranean basin according to the role of baroclinicity, diabatic process and orography (Fita et al, 2005). Severe weather is mainly associated with cyclones, but not necessarily with strong cyclones (Campins et al, 2005b), especially for heavy precipitation. Impact studies have concentrated on heavy rain (many authors), wind storms (for example Bora, Horvath et al, 2005), snowfalls (Pascual et. al, 2005). The impact in terms of heavy precipitation often depends on the presence of coastal mountains and of the relative position of cyclones in various areas (Ramis et al, 1994; Jansa et al, 2001).

- Among the many case studies, the Algerian-Balearic event ('superstorm') of Nov. 2001 received probably the widest attention for its exceptional intensity and disastrous consequences due to flash flood in Algiers and wind storm in the Balears. (Davolio and Buzzi, 2004; Arreola et al., 2004; Romero et al, 2004; Lambert et al, 2004; Tripoli et al, 2005; Smith et al, 2005). Regarding others strong historical events of Mediterranean cyclones, numerical reconstruction using ERA and NCEP reanalyses have been made (e.g. cyclone of Dec. 1979, Homar, 2002; Nov. 1966 event over Italy - Grossi et al, 2004; De Zolt et al, 2005).

3. **Methods and techniques used to draw the hypotheses/conclusions.**

- The most commonly applied diagnostic tools are: vorticity balance, PV dynamics, energetics (Prezerakos and Michaelides, 1989; Michaelides, 1992; Michaelides et al, 1999), Q-vectors; atmospheric trajectories with clustering to define "air mass" and water vapour paths (Bertò et al, 2004); application of the factor separation technique (Stein and Alpert, 1993; Alpert et al, 1995; Homar et al. 1999; Romero et al. 2000; Krichak and Alpert, 2002; Genoves and Jansa, 2002; Horvath et al, 2005).
- PV inversion technique: despite constant improvements of the observing and data assimilation systems, uncertainties in the initial conditions are still a major source of error for the numerical weather prediction models. When incipient analysis errors can be detected (e.g with the help of satellite imagery), alternative initial conditions (and subsequent forecast) can be obtained by editing and inverting the PV field. This method was used to investigate the sensitivity of the 2001 Algerian cyclone to sub-synoptic upper-level PV features (Lambert et al., 2004; see also Romero, 2001, and Romero et al, 2005).
- Statistical approach has been applied to relate cyclone characteristics and severe weather (e.g. heavy precipitation), using, for example, analogues and cluster analysis. Using a satellite-based climatology, Mediterranean cloud systems and troughs were automatically detected with the retrievals of the cloud top pressure and the temperature of the lower stratosphere. When focused on the south Alpine region, it is found that the most intense precipitating cases are associated with the most pronounced and elongated upper-level anomalies (Chaboureau and Claud, 2005). Martius et al (2005) obtained similar results through a climatological analysis between PV streamers and heavy precipitation events.

4. Numerical modelling (including analysis/assimilation) aspects.

- In order to correctly simulate and predict cyclogenesis and cyclone life cycle, it is necessary to “capture” (through assimilation and modelling) a range of scale and scale interaction processes from the quasi-planetary scale (Atlantic storm-track, mid-latitude and sub-tropical circulation) down to local scale (convection and orographic forcing). (Corazza et al, 2003; Richard et al, 2003; Accadia et al, 2005; Zampieri et al, 2005; Casaioli et al, 2005). Interaction between Atlantic tropical cyclones and mid-latitude flow may decrease the predictability of Rossby wave packets propagating towards the Mediterranean (Enamoto et al, 2005). Despite the complexity of many interacting factors, some high impact events over the Mediterranean like heavy precipitation south of the Alps may have a better predictability than average conditions through their link with large anomalies in the upper-level flow (Grazzini, 2005).
- In recent years there has been a gradual migration of models to finer scales, in scientific and especially operational models, with new approach to the implementation of operational convective-resolving models. Several case studies have been used to test models: MAP cases (e.g. IOP2b); Monserrat 2000 (Mariani et al, 2005; Zampieri and Buzzi, 2005); Aude 1999 (Ducrocq et al, 2002); Gard 2002 (Ducrocq et al, 2003; Llasat et al, 2003; Anquetin et al, 2005; Chancibault et al, 2005; Rigo and Llasat, 2005); Algiers 2001 'superstorm'; Israel heavy rain in 2001 (Krichack et al, 2004), and others (Federico et al, 2005), etc... Various problems have still to be solved, due to numerical resolution and associated numerical diffusion (Cheruy et al, 2004), to parameterization of turbulence, microphysics etc..
- Model verification, validation and intercomparison poses several issues and requires techniques that have been discussed and applied in the context of Mediterranean meteorology (Cardinali et al, 1998; Calenda et al, 2000; Kotroni and Lagouvardos, 2001; Picornell et al, 2002; Lagouvardos et al, 2003; Accadia et al, 2003; Accadia et al, 2005; Anquetin et al, 2005; Mariani et al, 2005; Anquetin et al, 2005; Yates et al, 2005), including the development of the model-to-observation approach (Caumont et al, 2005; Brenot et al, 2005).
- Models and model-related techniques have been applied to conduct sensitivity studies based on PV inversion and on the application of adjoint techniques. Identification of sensitive areas for Mediterranean cyclones has been done for a number of Mediterranean cyclones (Romero et al. 2001, Homar and Stensrud 2004, Romero et al., 2005; Homar et al, 2005). Such areas are spread both inside and outside the Mediterranean, depending on the large scale pattern, flow direction and on the time lag considered. Poorly covered areas (mid-low latitude eastern Atlantic and north Africa) are often sensitive areas. A climatological study has also confirmed that the lower troposphere of a region covering most of north Africa (up to 20 N) is potentially very sensitive for the quality of the short-term forecast over southern Europe (Marseille, 2001).
- Assimilation studies have been applied to determine the sensitivity of forecasts Mediterranean cyclones and associated severe weather to specific data assimilation, as for example assimilation of humidity (Ducrocq et al, 2002); assimilation of precipitation (for the Algiers storm: Davolio and Buzzi, 2004); humidity adjustment (application in 16 cases over Eastern Mediterranean, Lagouvardos and Kotroni, 2005); assimilation of space observations (Lanciani, 2004; Mugnai, 2005, others...).

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